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# A geostatistical approach to the simulation of stacked channels

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#### ABSTRACT

Turbiditic channels evolve continuously in relation to erosion-deposition events. They are often gathered into complexes and display various stacking patterns. These patterns have a direct impact on the connectivity of sand-rich deposits. Being able to reproduce them in stochastic simulations is thus of significant importance. We propose a geometrical and descriptive approach to stochastically control the channel stacking patterns. This approach relies on the simulation of an initial channel using a Lindenmayer system. This system migrates proportionally to a migration factor through either a forward or a backward migration process. The migration factor is simulated using a sequential Gaussian simulation or a multiple-point simulation. Avulsions are performed using a Lindenmayer system, similarly to the initial channel simulation. This method makes it possible to control the connectivity between the channels by adjusting the geometry of the migrating areas. It furnishes encouraging results with both forward and backward migration processes, even if some aspects such as data conditioning still need to be explored. © 2017 Elsevier Ltd. All rights reserved.

#### 1. Introduction

Heterogenities within turbiditic channel deposits can have a dramatic impact on fluid flow and reservoir production (e.g., Gainski et al., 2010). Mud-rich deposits such as margin drapes or slumps can obstruct fluid circulation and compartmentalize the reservoir depending on the stacking pattern (Labourdette et al., 2006), i.e., how channels position themselves in relation to each others. Significant changes in the stacking pattern can be observed even over short distances (Mayall and O'Byrne, 2002). This represents a major source of uncertainty regarding the connectivity, and modeling the stacking can help to assess this uncertainty.

This stacking results from two main processes: channel migration and avulsion. Migration occurs either through the gradual erosion and accretion of sediments along the channel margins, called continuous migration (e.g., Abreu et al., 2003; Arnott, 2007; Nakajima et al., 2009), or through the incision and filling of a new channel, sometimes with a significant distance between the channels, called discrete or abrupt migration (e.g., Abreu et al., 2003; Deptuck et al., 2003; Maier et al., 2012). Four patterns stand out (Fig. 1):

- A lateral channel bend migration or swing, which shifts the bend laterally and increases the channel sinuosity (Peakall et al., 2000; Posamentier, 2003).
- A downsystem channel bend migration or sweep, which shifts the bend downward (Peakall et al., 2000; Posamentier, 2003).
- A channel bend retro-migration, which decreases the channel sinuosity (Nakajima et al., 2009).
- A vertical channel migration or aggradation, which shifts the channel upward (Peakall et al., 2000).

Avulsion occurs when the density currents exceed the channel capacity to contain them: the flow leaves the channel and forms a new course.

Simulating migration and avulsion is a central research subject to better model the channel stacking. In fluvial systems, the most widespread methods are two-dimensional physical simulations (Lopez, 2003; Pyrcz et al., 2009). They link the migration to the asymmetry in the flow field induced by the channel curvature, which is responsible for bank erosion (Ikeda et al., 1981). These two-dimensional physical methods have been applied (McHargue et al., 2011) and adapted (Imran et al., 1999) to turbiditic environments. But the physical processes behind submarine channels remain controversial. The main controversy concerns the rotation direction of the secondary flow and its controlling factors (e.g., Corney et al., 2006, 2008; Imran et al., 2008), which constrain channel migration. Dorrell et al. (2013) argue that two-dimensional



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**Fig. 1.** Example of channel migration patterns interpreted on seismic data from the Benin-major channel-belt, near the Niger Delta (modified from Deptuck et al. (2003)).

physical models are not accurate enough to capture the full threedimensional structure of the flow field. This lack of accuracy was also pointed out in fluvial settings, especially with simplified physical models (Camporeale et al., 2007). More complex twodimensional models or three-dimensional models call for more parameters and a bigger computational effort, and their validity remains questioned (e.g., Sumner et al., 2014). Thus, their convenience in a stochastic framework is doubtful.

Another approach proposes to only reproduce some stratigraphic rules, for instance mimicking the migration without actually simulating the physical processes (Pyrcz et al., 2015). Viseur (2001) and Ruiu et al. (2015) defined migration vectors from a weighted linear combination of vectors for lateral migration, downsystem migration, and bend rotation. Teles et al. (1998), Labourdette (2008) and Labourdette and Bez (2010) went one step further by defining empirical laws controlling the spatial structure of the migration from modern channels or channels interpreted on seismic data. All such methods derive from objectbased approaches, which simulate a channel object. On the other hand, cell-based approaches (e.g., Deutsch and Journel, 1992; Mariethoz et al., 2010) paint the channels inside a grid based on a prior model. This prior model describes spatial structures and their relationships. Such methods can simulate almost any structure with few parameters. However, they have difficulties reproducing continuous channelized bodies and are not designed to model channel migration. Here we propose a different approach to channel migration, combining object- and pixel-based approaches as done with other geological structures (e.g., Caumon et al., 2007; Zhang et al., 2009; Rongier et al., 2014).

The proposed channel migration method uses a geostatistical simulation to reproduce the spatial structure resulting from the physical processes rather than modeling the physical processes themselves. We stochastically simulate the spatial evolution of a channel from one stage to the next using either a sequential Gaussian simulation (SGS) or a multiple-point simulation (MPS) method (section 2). Such a descriptive approach avoids the use of physical models that can be difficult to parameterize. After describing its capabilities (section 3), the method is applied to a synthetic case of confined turbiditic channels (section 4). This case includes a comparison of the connectivity from migrated channels and from randomly implanted channels. Finally, we discuss those results along with some perspectives for the method (section 5).

#### 2. Stochastic simulation of channel evolution

We divide channel migration into two elements:

- The horizontal component (hereafter referred to as migration), which includes the lateral, downsystem, and retro-migrations.
- The vertical component (hereafter referred to as aggradation), which includes the vertical migration.

(Labourdette, 2008) proposed to initiate the process from the last channel of a system, so from the youngest channel, and migrate backward. Indeed, this last channel is often observable on seismic data due to its argileous fill. Then the migration divides in two processes:

- A forward migration, which is the normal or classical migration. It starts from the oldest channel in geological time, which migrates to obtain the youngest channel.
- A backward migration, which is a reverse migration. It starts from the youngest channel in geological time, which migrates to obtain the oldest channel.

We propose a process to handle both forward and backward channel migrations.

#### 2.1. Channel initiation

The simulation calls for an already existing channel to initiate the evolution process. This initial channel can be interpreted from seismic data with a high enough resolution. Otherwise, it must be simulated. Here we use a formal grammar, the Lindenmayer system (L-system) (Lindenmayer, 1968), for this simulation.

The L-system rewrites an initial string using rules, which replace a set of letters by another one. The resulting string is then interpreted into an object. Rongier (2016) defined some rules in a framework able to stochastically simulate a channel from a L-system. These rules tie channel bends together, controlling the bend morphology and the orientation change between each bend. It results in a channel centerline, i.e., a set of locations through which the channel passes. Non-Uniform Rational B-Spline (NURBS) surfaces dress the L-system to obtain the final channel shape (Ruiu et al., 2015).

This method simulates various meandering patterns, from straight to highly sinuous channels. It is suitable for both forward migration, which classically requires starting with a quite straight channel, and backward migration, which requires an initial channel with a high sinuosity.

#### 2.2. Channel migration

Channel migration is deeply linked to the channel curvature. Other elements have to be considered, such as soil properties or flow fluctuations. However, the physical processes behind bend evolution are complex and still not completely understood. Moreover, they differ from turbiditic to fluvial environments. This is why we propose to rely on a more descriptive approach based on geostatistics.

#### 2.2.1. General principle

In physical approaches, a migration factor is computed along the nodes of a channel centerline based on fluid flow equations. Then the nodes are moved based on that factor along the normal to the centerline.

We rely on a similar approach based on moving the nodes along the normal to the centerline to migrate a channel (Fig. 2). Here the Euclidean distance *d* of displacement for a node is the length of a displacement vector  $\vec{v}$ :

$$d = \|\overrightarrow{v}\| \tag{1}$$

The displacement vector divides into two components (Fig. 2):

- A vertical component for the aggradation. Aggradation is simply done by shifting the new channel vertically by an aggradation factor ε<sub>q</sub>, which is the same for all the channel nodes.
- A horizontal component defined by a migration factor ε<sub>m</sub> computed using a stochastic simulation method.

The stochastic simulation of the migration factor is done either



**Fig. 2.** Migration principle: the centerline nodes are moved along the vectors  $\vec{v_a}$  and  $\vec{v_m}$ .  $\vec{v_a}$  is the aggradation component along the vertical direction symbolized by the normalized vector  $\hat{z}$ . The aggradation factor  $e_a$  determines the vertical displacement.  $\vec{v_m}$  is the migration component along the normal direction to the centerline symbolized by the normalized vector  $\hat{n}$ . The migration factor  $e_m$  determines the horizontal displacement.  $\hat{s}$  is the normalized vector  $\hat{n}$ . The migration factor  $e_m$  determines the horizontal displacement.  $\hat{s}$  is the normalized vector along the streamwise direction.

with sequential Gaussian simulation (SGS) or multiple-point simulation (MPS). In both cases, the curvature becomes a secondary variable that influences the structuring of the migration factor. We detail some numerical aspects valid for both the SGS and MPS in the supplementary materials concerning:

- Curvature computation.
- Regridding to preserve a constant distance between the centerline nodes after migration (Fig. 2).
- Smoothing to eliminate undesired small-scale fluctuations of the migration factor, which can have a significant impact after some migration steps.

#### 2.2.2. Migration through sequential Gaussian simulation

The SGS simulates a migration factor value for each node of the centerline in a sequential way (e.g., Deutsch and Journel, 1992). Here we use an intrinsic collocated cokriging (Babak and Deutsch, 2009) to introduce the curvature as secondary variable:

- 1 A random path is defined to visit all the centerline nodes.
- 2 At a given node:
  - (a) If some nodes in a given neighborhood already have a value:
    (i) A kriging system determines the Gaussian complementary cumulative distribution function (ccdf) using the data, i.e., the nodes with a value given in input and the previously simulated nodes within the neighborhood.
    - (ii) A simulated value for the given node is drawn within the ccdf.
  - (b) Otherwise, the simulated value is drawn from an input distribution of migration factor.
- 3 Return to step 2 until all the nodes of the path have been visited.

The SGS requires the migration factor to be a Gaussian variable. If not, a normal score transform of the input distribution and of the data is introduced before step 1. A back transform is done at the end of the simulation process.

This simulation of the migration calls at least for four parameters (Fig. 3, see the supplementary materials for more details):

- Two distributions, one for the aggradation factor ε<sub>a</sub> and one for the migration factor ε<sub>m</sub>. They control the distance between two successive channels.
- A variogram range *r*. It controls the extension of the migrating area along a bend. The other variogram parameters get default values.
- A curvature weight γ<sub>c</sub>. It represents the correlation between the primary variable, i.e., the migration factor, and the secondary variable, i.e., the curvature. When this weight is positive, the channel tends to migrate, when it is negative, the channel tends to retro-migrate. Thus, simply by changing the curvature weight symbol the same workflow achieves both forward and backward migration processes.

#### 2.2.3. Migration through multiple-point simulation

Simulation methods such as the SGS rely on a histogram and a variogram inferred from the data. Thus, they only catch the oneand two-point statistics and miss all the higher-order statistics. But higher-order statistics are difficult if not impossible to infer from data. Multiple-point simulation (Guardiano and Srivastava, 1993) attempts to overcome such limitation by relying on an external representation of the structures of interest, the training image. Here the training image is a set of migrating channels, with the

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**Fig. 3.** Main parameters used for horizontal bend migration with SGS.  $e_m$  is the migration factor, *r* the variogram range and  $\gamma_c$  the curvature weight. The later perturbs the two other parameters by fitting more or less the migration spatial structure to the curvature spatial structure.

aggradation factor, migration factor and curvature values all known. Using MPS instead of SGS could lead to more realistic migrations by using real channels as training set.

The whole training set is not necessarily used to simulate the migration of a channel. The process relies on a training model, which can be (Fig. 4):

- The entire training set. In this case, each simulated migration step is influenced by all the migration steps within the training set.
- A single migration step within the training set:
  - -Drawn randomly among all the migration steps of the training set.

-That follows the migration order of the training set. In that case, each simulated migration step corresponds to a particular migration step within the training set. With that option, the number of migration steps in the training set limits the number of simulated migration steps.

The aggradation values are randomly drawn from the training set and attributed to all the nodes of the channel to migrate. The migration factor is simulated using the Direct Sampling method (DS) (Mariethoz et al., 2010):

1 A random path is defined to visit all the centerline nodes.

- 2 At a given node (Fig. 5):
  - (a) The *n* closest nodes with already a value form two data events  $N_x$ , one for the migration factor and one for the curvature.
  - (b) Those data events are searched within the training model: i A position is randomly chosen and the training model is then scanned linearly.
    - ii At a given node:

A Two distances  $d_{D,P}$  are computed between the current pattern  $N_y$  and the data event  $N_x$ , one for the migration factor and one for the curvature:

$$d_{D,P}(N_x, N_y) = \frac{1}{n} \sum_{i=1}^n \frac{|Z(x_i) - Z(y_i)|}{\max_{y \in TM} (Z(y)) - \min_{y \in TM} (Z(y))}$$
(2)

with  $d_{D,P} \in [0, 1]$ , *n* the number of nodes in the data event, *Z* the compared property, *x* a node in the data event, *y* a node in the training set pattern and *TM* is the training model.

- B If the distances are both the lowest encountered, the value of the central node associated to the pattern is saved.
- C If the distances are both lower than given thresholds, the scan stops.

(c) The saved value becomes the simulated value for this node.3 Return to step 2 until all the nodes of the path have been visited.

This method has the advantage of easily handling continuous properties and secondary data, here the curvature. The curvature ensures the link between the spatial variations of the migration factor in the training model and in the simulation.

Besides the training model, this simulation of the migration calls for the classical DS parameters (see the supplementary materials for more details):

- The maximal number of nodes in a data event.
- The maximal portion of the training model to scan.
- A threshold for the migration factor.
- A threshold for the curvature.

#### 2.3. Neck cut-off determination

As the channel sinuosity increases, the two extremities of a bend come closer one to the other until the flow bypasses the bend. This is a neck cut-off, leading to the abandonment of the bypassed bend.

As done in physical simulations (e.g., Howard, 1992; Camporeale et al., 2005; Schwenk et al., 2015), neck cut-offs are simply identified when two non-successive nodes of the centerline are closer than a given threshold. The lowest threshold is the channel width, as the margins of the bend come in contact. However, this threshold is quite restrictive and not so realistic (Camporeale et al., 2005). Here the threshold is set to 1.2 times the maximal channel width.

The search for cut-off starts upstream and continues to the most downstream part of the channel. The distance between a given node and another non-successive node of the centerline is compared with the threshold. When the distance is lower, these two nodes and all the nodes in-between are suppressed. The cutting path is then symbolized by two nodes. A new node is added along that path (Schwenk et al., 2015), using a cubic spline interpolation. This method of neck cut-off determination is simple but rather time-consuming. More efficient methods exist to reduce the computation time (e.g., Camporeale et al., 2005; Schwenk et al., 2015).

For now only the forward migration process handles the formation of neck cut-offs. Indeed, the cut-offs appear naturally as the sinuosity increases. In the backward process, the sinuosity decreases: introducing neck cut-offs calls for a different method, which is a perspective of this work.

#### 2.4. Avulsion

Avulsion is a key event widely observed in both fluvial and turbiditic systems. When an avulsion occurs, the channel is



Fig. 4. Framework for the simulation of the migration factor  $e_m$  with MPS.

abruptly abandoned at a given location (Fig. 6).

Upstream, the flow remains in the old channel, whereas downstream a new channel is formed. However, its triggering conditions remain poorly understood due to the complexity of this process. Avulsion is often statistically handled in simulation methods: a probability of avulsion controls the development of a new channel. This process can be influenced by the curvature, as a high curvature tends to favor an avulsion.

The approach for global avulsion is similar to that defined by Pyrcz et al. (2009). The avulsion starts by computing the sum of the curvatures at each channel section. A threshold is randomly drawn between zero and the sum of curvatures. The channel is then scanned from its most upstream part to the downstream part. At each centerline node, the curvature is subtracted from the threshold. A section initiates an avulsion or not depending on two factors:

- An input probability of avulsion.
- The random curvature threshold, which should be lower than the section curvature to trigger an avulsion.

Thus, the avulsion initiation at a given section is a probabilistic choice influenced by the curvature at the section location.

Then, the upstream part of the channel is isolated. It becomes the initial string to simulate the new post-avulsion channel with a L-system (Fig. 6). This channel is based on the same parameters than the initial channel, but different parameter values may be used. A repulsion constraint (Rongier, 2016) with the pre-avulsion channel can be set to avoid intersections between the two channels.

#### 3. Applications

The method was implemented in C++ in the Gocad plug-in ConnectO. The channel envelopes based on NURBS were implemented by Jérémy Ruiu in the Gocad plug-in GoNURBS (Ruiu et al., 2015).

The method was used to simulate two realizations: one following a forward migration process and one following a backward migration process (Fig. 7).

Both processes are able to reproduce various migration patterns, from lateral to downsystem migration, and even areas of retromigration (Fig. 8). Some bends also evolve to complex bends constituted by several bends: this leads to the formation of new meanders. These synthetic cases have been developed without any conditioning data. The range is chosen similar to the bend length. The other variogram parameters are those predefined (see the supplementary materials). The curvature weight is kept high, giving a dominant lateral migration. The forward process could keep migrating over more steps, with neck cut-offs keeping the channels within a restrained area. The backward process does not migrate much after a few steps, when the channel starts to miss significant bends. An avulsion and sometimes an abrupt migration can reestablish some migration.



**Fig. 5.** DS simulation principle: the neighboring configuration around the node to simulate, the data event, is sought within the similar configurations in the training model, the patterns. Here the neighboring configurations contains the four nodes with a known value that are the closest to the node to simulate. When the distance  $D_{DP}$  between the data event and a pattern is lower than a given threshold  $d_t$ , the process stops and the node to simulate gets the value at the same location within the pattern.

Abrupt migrations are handled by introducing a second set of migration parameters. The channel centerline is scanned upstream to downstream. A probability of abrupt migration defines if an abrupt migration occurs. The appearance of an abrupt migration is also weighted by the channel curvature. When an abrupt migration occurs, an abrupt migration length is drawn from an input distribution. All the nodes along the drawn length migrate following the second set of migration parameters. Such abrupt migration process can introduce a spatial discontinuity from the previous channel (Fig. 8).

Avulsion momentarily stops the migration process, which restarts from the new channel. The continuity between the upstream part to the avulsion location and the newly simulated channel is finely preserved (Fig. 8). The use of the curvature to weight the abrupt migration and avulsion process tends to decrease their emergence in the backward process. This is due to the sinuosity decrease induced by such process. In this case, higher probabilities are used.

Neck cut-offs naturally appear during the forward process as the sinuosity increases (Fig. 8, forward migration). The proposed backward process is unable to generate cut-offs. As the migration advances, the channel just gets straighter. It does not evolve to a complete straight line, but continuing the process does not lead to an increase of the sinuosity: the channel remains in a steady-state.

To test the method with MPS, a training set was simulated with a SGS-based process (Fig. 9).

This training set has 9 migrating steps. Lateral migration

Current channel



Fig. 6. Principle of global avulsion based on L-system.

dominates the system, and several abrupt migrations perturb the channel stacking. Each migration step simulates the migration factor from the corresponding step in the training set, and not from the entire training set. The parameters for the MPS favor quality over speed, with the two thresholds at 0.01 and maximal scanned fraction of the training model of 0.75. But the simulated migration factors are noisy, because the process can not always find a pattern that meets the thresholds. We added two smoothing iterations (see the supplementary materials) at the end of each migration step to limit the noise.

In the end, the lateral migration is still dominant in the simulations. Abrupt migrations are also reproduced by the simulation process. But they tend to be less frequent. They also tend to be smaller, both in length and in migration factor, than in the training set. This comes from both the inability to find the right pattern in the training set and from the smoothing.

# 4. Analysis of the channel stacking impact on the static connectivity

This section aims at highlighting the impact of the migration



Fig. 7. Examples of realizations for both the forward and backward SGS migration processes SGS initiated by channels simulated with a L-system. The input parameters are given in the supplementary materials.

process on the simulated channel connectivity. To do so, we rely on a synthetic case study including realizations with different stacking

patterns.

#### Forward migration



Fig. 8. Enlargements on some areas of the channels on Fig. 7 illustrating different aspects of channel evolution reproduced by forward and backward migration simulations.

#### 4.1. Case study

The case study is inspired by turbiditic systems and their nested channelized bodies (e.g., Abreu et al., 2003; Mayall et al., 2006; Janocko et al., 2013). In such settings, channels migrate within and gradually fill a master channel – a large incision that confines

#### the channels:

• Lateral migration dominates the first phase of the filling. The channels migrate within the whole master channel width, with a low aggradation and some abrupt lateral migrations. Sand-rich





Fig. 9. Application of a forward migration process based on MPS to two channels generated with L-system. The input parameters are given in the supplementary materials.

channel deposits occupy the entire bottom of the master channel.

• Aggradation dominates the second phase of the filling. The lateral migration is less significant, and no abrupt migration arises. Sand-rich deposits occupy a limited area within the top of the master channel. The rest of the master channel is filled with inter-channel deposits, in particular inner levees whose development induces the limited lateral migration.

A hexahedral grid aligned along its margins represents the master channel (Fig. 10). Three sets of 100 realizations are simulated within this grid, with each realization containing 40 channels (see the supplementary materials for the input parameters).

The first two sets rely on a traditional object-based procedure: the channels are randomly placed inside the grid. Here, each channel is simulated using a L-system, similarly to the initial channel for the migration. L-systems condition data thanks to attractive and repulsive constraints (Rongier, 2016): in the proposed application, the master channel margins repulse the channels to keep them confined. The first set (Fig. 11, a) is limited to this setting: its channels are free to occupy the whole grid without any constraint on their relative location. These channels display a disorganized stacking. The second set (Fig. 11, b) further uses the L-system ability to condition data to reproduce the two-phase evolution of the channels. The random channel placement and the channel development are both influenced by a sand probability cube that defines the sand-rich deposit distribution inside the grid (Fig. 10). It influences the relative positions of the channels without directly controlling the channel relationships. These channels display a conditioned disorganized stacking.

The last set (Fig. 11, c) also attempts to reproduce the two-phase channel evolution, but without using the probability cube. Instead, it directly simulates this evolution with a forward SGS-based migration. The first 27 migration steps simulate a high lateral migration, some abrupt migrations and little aggradation. This first phase is initiated with a channel simulated with a L-system, whose initial position is randomly drawn at a fixed vertical coordinate along the bottom of the grid. The next 12 steps simulate a small lateral migration with a significant aggradation. This second phase is initiated with the last channel of the first phase. If a channel node



Fig. 10. Dataset of the application: a curvilinear grid representing a master channel with a sand probability cube.



Fig. 11. Examples of realizations corresponding to different approaches for channel simulation. Each realization contains 40 channels within a master channel.

should migrate outside the master channel, its migration factor value is decreased so that the channel remains within the master

channel, along its margin. In this set, the migration dictates the channel relationships. These channels have an organized stacking.

#### 4.2. Connectivity analysis principle

The connectivity analysis helps to compare realizations by focusing on the connectivity of the sedimentary deposits (Rongier et al., 2016). It relies on indicators based on the connected components of the different deposit types and their curve-skeletons (Table 1). These indicators give some information about the proportion of deposits, their connections or their shape. Here we only consider the channel deposits in the analysis, and not the inter-channel deposits within the master channel.

Dissimilarity values between the realizations facilitate the analysis. The dissimilarities compare and combine the indicators by means of a heterogeneous Euclidean/Jensen-Shannon metric. To

#### Table 1

Set of indicators and associated weights used for the case study. Indicator descriptions are in Rongier et al. (2016). Three other indicators exist but are nondiscriminant in this case, so not used: the *facies adjacency proportions*, because the realizations only contain two facies; the *unit connected component proportion*, because the rasterized objects do not lead to any connected component of one cell; the *traversing connected component proportion*, because all the channel objects go through the entire master channel and are all traversing.

Category	Indicator	Weight
Global indicators	Facies proportion	1
	Facies connection probability	1
	Connected component density	1
Shape indicators	Number of connected component cells	1
	Box ratio	1
	Faces/cells ratio	1
	Sphericity	1
Skeleton indicators	Node degree proportions	1
	Inverse branch tortuosity	1

analyze them, Rongier et al. (2016) proposed to use the Scaling by MAjorizing a COmplicated Function (SMACOF) (De Leeuw and Mair, 2009), a multidimensional scaling (MDS) method. The purpose is to represent the realizations as points, so that the distances between the points are as close as possible to the dissimilarities between the realizations. It implies that the MDS may paint an erroneous picture of the dissimilarities. The Shepard diagram and the scree plot help to assess the dissimilarity reproduction: the lower the stress on the scree plot and the higher the linear regression coefficient on the Shepard diagram, the better the representation is.

#### 4.3. Indicator results

The indicators and dissimilarities help to objectively analyze and compare the difference in connectivity between the three sets: organized stacking, conditioned disorganized stacking and disorganized stacking.

#### 4.3.1. Global analysis on the dissimilarity values

The multidimensional scaling plots the dissimilarities in a twodimensional representation (Fig. 12). The Shepard diagram and the scree plot show that two-dimensions are sufficient to represent the dissimilarities without significant bias. Three dimensions would have been a bit better, but more difficult to analyze.

The dissimilarities clearly divide the realizations in two groups. The first group contains all the 100 organized stacking realizations, 47 conditioned disorganized stacking realizations and 58 disorganized realizations. The realizations of the different sets do not mix much, with three sub-groups, one per realization set. The conditioned disorganized stacking realizations are closer to the organized stacking realizations than the disorganized stacking



Fig. 12. Multidimensional scaling representation comparing three sets of realizations with different methods and parameters. The identified realizations are shown in Fig. 13.

realizations. The second group contains 53 conditioned disorganized stacking realizations and 42 disorganized stacking realizations. Compared to the first group, the realizations are a bit more mixed, with a significant variability between the realizations.

Visually, the difference between the realizations of the different sets is quite clear (Fig. 13). However, looking at realizations from the same set but in different groups does not show any significant difference.

#### 4.3.2. Detailed analysis on the indicator values

Examining the indicators explains the separation into two groups (Fig. 14). The realizations of the first group all have a facies connection probability of one. These realizations have channel deposits that form a single connected component. The second group contains all the realizations with more than one connected component. This highlights the continuity of the migration process: having a complete non-connection between two successive channels requires an avulsion. Migration makes it easier to control the channel connectivity.

Within the first group, the disorganized stacking realizations are clearly different from the other realizations. This appears on the facies proportion and on the average number of component cells. With the same number of channels, the connected components of these realizations are larger than those of the other sets due to the disorganized stacking. On the other hand, the conditioned disorganized stacking and organized stacking realizations have similar facies proportions and average numbers of component cells. Their difference appears on the other indicators, such as the average faces/cells ratio or the average sphericity: even if the channels of the two sets occupy similar volumes within the grid, their shapes are different. The low faces/cells ratio of the organized stacking realizations highlights their structure: the channels are significantly stacked over long distances, which decreases more the number of faces of the components than their number of cells. The average sphericity of these realizations is higher than that of the conditioned disorganized realizations. This comes from their respect of the channel evolution: they occupy the whole width of the master channel bottom, and vertically they evolve to the top of the master channel. This also comes from the management of the channel margins: the migration is simply blocked by the channel margins, which is less constraining than the margin repulsion, especially at the bottom of the grid.

The difference between the realization sets within the first group is also visible on the skeletons (Fig. 15). The disorganized stacking realizations have higher proportions for the node degrees larger than 3 compared to the realizations of the other sets. This highlights channels that locally cross each other but are globally disconnected. This tends to generate many small branches all along the skeleton, with many loops (Figs. 16 and 13). The difference between the conditioned disorganized stacking and the organized stacking realizations of the first group is less significant than on the other indicators. However, the conditioned disorganized stacking realizations have higher proportions for the node degrees larger than 3. Again, this highlights the tendency of their channels to cross each other instead of stacking on each other (Fig. 16). This is visually striking on the skeletons (Fig. 13): the conditioned disorganized stacking realizations have many small branches forming loops, similarly to the disorganized stacking realizations. The organized stacking realizations have fewer small branches. The small branches also tend to be straight, with an inverse tortuosity close to one. From this perspective, the evolution from the disorganized stacking to the conditioned disorganized stacking and the organized stacking is clear on the average inverse tortuosity: as the stacking increases, the inverse tortuosity decreases due to less straight branches within small loops.

Thus, the difference in stacking directly impacts the shape of the connected components and their connectivity. Adding a sand probability cube helps to control the connectivity between the channel deposits. But the resulting channels do not stack as clearly as with the migration process, which prevents non-connections if required.

#### 5. Discussion and perspectives

The previous applications highlight the relevance of the migration approach. The following section discusses some aspects of the method.

#### 5.1. About the migration pattern simulation

As defined, the process based on SGS leads to a dominant lateral migration through the influence of the curvature. Using a low curvature weight leads to the random emergence of other patterns. Some asymmetric bends can also randomly appear, with a random orientation of their asymmetry. It is not possible to choose another dominant migration pattern, such as a downsystem migration. This is not an issue for turbiditic channels, which tend to have little downsystem migration (e.g., Nakajima et al., 2009). However, if another dominant pattern is required, the method must be adapted. A simple solution is to modify the vector of migration by adding a downsystem component, such as done by Teles et al. (1998) or Viseur (2001). Another solution is to change the secondary data influencing the migration. The curvature could be modified, or a different property could have to be used. From this point of view, the MPS has the advantage that the training set controls the migration pattern: any pattern on the training set can appear in the realizations.

Globally, the SGS remains able to reproduce various migration patterns. Most of the time, no smoothing is required. The MPS still needs more work to improve the reproduction of the migration pattern from the training set. It usually calls for a training set larger than the realizations to increase the repeatability of the patterns and improve the realization quality. From this point of view, the training set used in Fig. 9 is not optimal, because the channels have roughly the same length than those in the realizations. It leads to small-scale perturbations in the migration factor that deform the meanders after a number of migration steps. The smoothing helps to limit these perturbations and to obtain more realistic results.

The bends should also be compared with their counterparts from real cases to assess the method's ability to simulate a realistic migration. MPS should perform better when using a real case as training set, but this needs to be further tested. Statistics such as those of Howard and Hemberger (1991) can be used to compare the migrating channels. But they are not directly defined to analyze the migration. Histogram and variogram of the migration factor can give a first insight, but further indicators should be developed to objectively analyze and compare migration patterns.

A comparison could be done with physical simulation methods, especially the stochastic ones (Lopez, 2003; Pyrcz et al., 2009). The main uncertainty comes from the ability of the physical model to explore all the possible migration patterns. For instance, our method is able to simulate some retro-migrating areas. These areas form outer-bank bars, which are potential reservoir areas (Nakajima et al., 2009). Such bars have no equivalent in fluvial processes, whereas all the physical methods for the migration are developed for the fluvial environment. Thus, they may not be able to develop such migration patterns. Comparing the results of the forward and the backward migration processes would be also an interesting development.



Fig. 13. Realizations and their skeletons for each set within the two groups separated by the dissimilarities. Each realization is the closest to the mean MDS point of its set and group (see Fig. 12).



**Fig. 14.** Box-plots comparing the range of indicators – except the node degree proportions – computed on three sets of realizations with different methods and parameters. *OS1.* Organized stacking realizations within the group 1; *CDS1 and CDS2.* Conditioned disorganized stacking realizations within the groups 1 and 2; *DS1 and DS2.* Disorganized stacking realizations within the groups 1 and 2.

#### 5.2. About the discrete process simulation

For the SGS process, abrupt migrations are introduced by a second set of migration parameters. Discontinuities can then develop between channels. However, they tend to follow a similar migration pattern. Abrupt migrations and even local avulsions could also be simulated using L-systems. The initiation process would be the same as for avulsion. The newly simulated part would be attracted to a downstream location of the initial channel. This process is similar to the one developed by Anquez et al. (2015) to simulate anastomotic karst networks. It would possibly simulate bends completely independent from the previous channel.

MPS has shown its ability to reproduce abrupt migrations from the training set. Again, it makes the simulation easier once a training set is available. The only drawback is when avulsions or cut-offs are present in the training set. For now they are not handled, but simulating them based on their appearance in the training set could improve the process.

The appearance of neck cut-offs is not a problem in a forward process with SGS. With SGS in a backward process, neck cut-offs should be simulated during the process. This would let the migration continue over any number of migration steps.

#### 5.3. About the parameterization

Using SGS does not call for an intensive parameterization, with only four parameters required for a simulation. The aggradation and migration factors are directly related to the vertical and horizontal distances between two successive channels. They are thus pretty easy to define. The curvature weight is a bit harder to infer. A weight of 1 gives a significant influence to the curvature. By default, a weight around 0.8 gives a dominant lateral migration but lets other migration patterns appear. The variogram range should be close to the desired length of the bends that form during migration.



**Fig. 15.** Mean node degree proportions of the levee skeletons for each set and group. The error bars display the minimum and maximum proportions. The first 1 node degree corresponds to the nodes of degree one along a grid border. The second 1 node degree corresponds to the nodes of degree one inside the grid.

No use of conditioning data has been done yet to find the parameter values. This could be done from channels interpreted on seismic data. Even if all the channels are often not discernible, some of them could inform about the possible values for the factor distributions, the variogram parameters and even the curvature weight by comparing the channel curvature with the migration distance. Analogs from outcrops or seismic data of similar settings can also help to define these values.

One possibility to reduce the number of parameters is to use the bend length as range. The range then varies following the channel bends and the migration step. However, the channels often tend to develop small-scale variations that perturb the bend identification and thus the bend length computation. No significant migration can be obtained with such parameterization. One possibility is to smooth the simulated migration factors or the bend lengths. A better solution would be to better identify the bends and avoid small-scale variability. The work of O'Neill and Abrahams (1986) for instance could be a first lead.

Compared to physical methods (e.g., Ikeda et al., 1981; Parker

et al., 2011; Lopez, 2003), the parameterization is far simpler when the purpose is to model the current aspect of the geology. This requires working on old channels that have been deformed. Thus, the physical parameters that lead to the channel formation are difficult if not impossible to infer. Pyrcz et al. (2009) manage to reduce the number of parameters to a single maximum distance to reach by a standardization process. The impact of such standardization on the migration process and thus on the stacking patterns is not discussed. This parameterization is easier to infer, but less flexible if the migration patterns are not those desired. The parameters used here give a finer control to the user on the migration patterns. Furthermore, they are mainly descriptive and can be inferred from the available data.

The more processes are introduced, e.g., abrupt migration, the heavier the parameterization tends to be. The MPS approach is then pretty useful. It requires few parameters that are more related to the ratio between the simulation quality and the simulation speed. The training set dictates the geological considerations, such as the presence of abrupt migrations or the dominant migration patterns. The main issue is to find a training set. The most interesting option is to find one from an analog, either seismic data such as done by Labourdette (2008) or possibly an outcrop. Satellite images are also interesting sources of training sets in fluvial settings.

#### 5.4. About small-scale variability and smoothing

Realization statistics often fluctuate around those of the prior model (e.g., Deutsch and Journel, 1992). This can lead to some noise or short-scale variability. Too high MPS thresholds can also lead to a higher small-scale variability than within the training model. Noise or short-scale variability can form inflexions. Such inflexions prevent from using the bend length as range for the SGS, as discussed in section 5.3. They also tend to grow during the migration, forming new bends at a smaller-scale than initially desired.

Smoothing the migration factor controls the small-scale variability by eliminating its influence on the migration. However, the smoothing impact is quite significant, as discussed by Crosato (2007) when smoothing the curvature. Four to five smoothing steps can be enough to completely modify the migration structure. It should then be used carefully. Another option would be to postprocess the realizations to improve the reproduction of the prior model. Simulated annealing, for instance, makes it possible to better reproduce the histogram and variogram through the minimization of an objective function (e.g., Deutsch and Journel, 1992).

#### 5.5. About the usefulness of the migration process

The comparison with randomly placed channels highlights the difference of static connectivity. Simulating the migration gives more control on the stacking pattern. This is especially useful due to the significant influence of the stacking pattern on the connectivity. Influencing the channel locations by a probability cube reduces the gap with the migration results. But the difference in connectivity remains significant.

The analysis of the connectivity could be further developed by introducing the channel fill. In particular mud drapes have a significant impact on the connectivity. And in such case controlling the stacking pattern is even more important.

#### 5.6. About the simulation process with migration

L-system are interesting to simulate the initial channel, especially for their ability to develop channels with different sinuosities. In the MPS case, methods such as that of Mariethoz et al. (2014) could also be interesting. They simulate channel centerlines based



Fig. 16. Two migrating channels with two local abrupt migrations and associated skeletons. An organized stacking of the two channels results in a single branch on the skeleton. The areas of abrupt migration, where the channels are not stacked anymore, result in a loop on the skeleton.

on MPS in a process similar to that used for migration. The initial channel could then be simulated based on the first channel of the training set.

Both SGS and MPS are able to simulate a forward or a backward migration. This backward process is particularly useful, as the last channel of a migrating sequence is far more often interpretable on seismic data than the first one (Labourdette, 2008). This allows initiating the process from the real data, instead starting from an unknown state and trying to condition the process to the last channel.

For now the channel width and thickness are simulated at the end of each migration step. As the width has a particular impact on the migration, it could be interesting to simulate them earlier. With MPS, the channel width and thickness could also be simulated from the training set instead of using SGS. Other geological elements could be integrated, such as the channel fill (e.g., Labourdette, 2007; Alpak et al., 2013). This is especially important due to its impact on the connectivity. Levees also need to be introduced (e.g., Pyrcz et al., 2009; Ruiu et al., 2015). When the channels migrate within a confinement such as a canyon, they can erode that confinement. Thus, they modify the confinement morphology, which should be taken into account.

#### 5.7. About data conditioning

Data conditioning of the migration process has not been explored yet. Both SGS and MPS can be used for data conditioning. If a datum is within a conditioning distance from the current channel, the migration process can be conditioned to that information. The conditioning distance corresponds to the area in which a channel node can migrate. This area is determined by the maximal migration and aggradation factors. To preserve the conditioning, smoothing cannot be performed at the data locations.

However, the process is more difficult when the data are outside the conditioning distance. One solution is to introduce a constraint that attracts the migrating channel to the data, similarly to the Lsystem conditioning (Rongier, 2016) or to the conditioning of process-based methods (Lopez, 2003; Pyrcz and Deutsch, 2005). This implies adjusting the appearance of discrete migrations and avulsions depending on the data and their location. Conditioning to a sand probability cube is also problematic, especially for handling avulsions. This may require identifying the large-scale trends within the cube. It is also important to note that the overall methodology requires an important work of data interpretation and sorting to possibly pre-attribute them to each migrating system.

#### 6. Conclusions

This work provides a basis for a more descriptive approach to channel migration that focuses on the spatial structure of the migration. The same approach stochastically simulates either forward or backward channel migration, starting with an initial channel simulated by L-system or interpreted on seismic data. The migration process is based on simulating a migration factor using sequential Gaussian simulation or multiple-point simulation with the curvature as secondary data. Four parameters are required by the SGS approach to adjust the migration patterns. The MPS approach calls for four parameters related to the simulation speed and quality and a training set that controls the migration patterns. Avulsion is performed by L-system simulation, as for the initial channel.

The first results are encouraging: they show a significant difference in connectivity from a process with no direct control on the channel stacking. Further work is required on some points, such as using the bend length as variogram range. Both SGS and MPS offer some conditioning ability, but only if the data are close to the channel. Data management at further distance could be done with attractive constraints, as done for initial channel conditioning (Rongier, 2016) or with physical methods (e.g., Lopez, 2003; Pyrcz and Deutsch, 2005). Neck cut-offs remain to be introduced in the backward process with SGS. The training set required by MPS could be better used to take into account cut-offs and avulsions. The channel fill should also be simulated to better assess the impact on the static connectivity. The method was developed for the simulation of turbiditic channels, but it could also be applied to fluvial systems.

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#### Appendix A. Supplementary data

Supplementary data related to this article can be found at http:// dx.doi.org/10.1016/j.marpetgeo.2017.01.027.

#### References

- Abreu, V., Sullivan, M., Pirmez, C., Mohrig, D., 2003. Lateral accretion packages (LAPs): an important reservoir element in deep water sinuous channels. Mar. Pet. Geol. 20, 631–648. http://dx.doi.org/10.1016/j.marpetgeo.2003.08.003.
- Alpak, F.O., Barton, M.D., Naruk, S.J., 2013. The impact of fine-scale turbidite channel architecture on deep-water reservoir performance. AAPG Bull. 97, 251-284. http://dx.doi.org/10.1306/04021211067.
- Anquez, P., Rongier, G., Collon, P., 2015. Stochastic simulations of karst networks with Lindenmayer systems. In: Proc. 35th Gocad Meeting. France, Nancy.
- Arnott, R.W.C., 2007. Stratal architecture and origin of lateral accretion deposits (LADs) and conterminuous inner-bank levee deposits in a base-of-slope sinuous channel, lower Isaac Formation (Neoproterozoic), East-Central British Columbia, Canada. Mar. Pet. Geol. 24, 515-528. http://dx.doi.org/10.1016/ j.marpetgeo.2007.01.006.
- Babak, O., Deutsch, C.V., 2009. An intrinsic model of coregionalization that solves variance inflation in collocated cokriging. Comput. Geosci. 35, 603-614. http:// dx.doi.org/10.1016/j.cageo.2008.02.025
- Camporeale, C., Perona, P., Porporato, A., Ridolfi, L., 2005. On the long-term behavior of meandering rivers. Water Resour. Res. 41, W12403. http://dx.doi.org/10.1029/ 2005WR004109.
- Camporeale, C., Perona, P., Porporato, A., Ridolfi, L., 2007. Hierarchy of models for meandering rivers and related morphodynamic processes. Rev. Geophys. 45, RG1001. http://dx.doi.org/10.1029/2005RG000185.
- Caumon, G., Tertois, A.L., Zhang, L., 2007. Elements for stochastic structural perturbation of stratigraphic models. In: Proceedings of Petroleum Geostatistics. European Association of Geoscientists & Engineers. http://dx.doi.org/ 10.3997/2214-4609.201403041.
- Corney, R.K., Peakall, J., Parsons, D.R., Elliott, L., Best, J.L., Thomas, R.E., Keevil, G.M., Ingham, D.B., Amos, K.J., 2008. Reply to discussion of Imran et al. on "The orientation of helical flow in curved channels" by Corney, et al. Sedimentology vol. 53, 249-257. http://dx.doi.org/10.1111/j.1365-3091.2007.00925.x. Sedimentology 55, 241-247.
- Corney, R.K.T., Peakall, J., Parsons, D.R., Elliott, L., Amos, K.J., Best, J.L., Keevil, G.M., Ingham, D.B., 2006, The orientation of helical flow in curved channels, Sedimentology 53, 249-257. http://dx.doi.org/10.1111/j.1365-3091.2006.00771.x.
- Crosato, A., 2007. Effects of smoothing and regridding in numerical meander migration models. Water Resour. Res. 43, W01401. http://dx.doi.org/10.1029/ 2006WR005087
- De Leeuw, J., Mair, P., 2009. Multidimensional scaling using majorization: SMACOF in R. I. Stat. Softw. 31, 1-30.
- Deptuck, M.E., Steffens, G.S., Barton, M., Pirmez, C., 2003. Architecture and evolution of upper fan channel-belts on the Niger Delta slope and in the Arabian Sea. Mar. Pet. Geol. 20, 649–676. http://dx.doi.org/10.1016/j.marpetgeo.2003.01.004. Deutsch, C.V., Journel, A.G., 1992. GSLIB: Geostatistical Software Library and User's
- Guide. Oxford University Press, New York.
- Dorrell, R.M., Darby, S.E., Peakall, J., Sumner, E.J., Parsons, D.R., Wynn, R.B., 2013. Superelevation and overspill control secondary flow dynamics in submarine channels. J. Geophys. Res. Ocean. 118, 3895–3915. http://dx.doi.org/10.1002/ igrc.20277.
- Gainski, M., MacGregor, A.G., Freeman, P.J., Nieuwland, H.F., 2010. Turbidite Reservoir Compartmentalization and Well Targeting with 4D Seismic and Production Data: Schiehallion Field, vol. 347. UK. Geological Society, London, pp. 89-102. http://dx.doi.org/10.1144/SP347.7. Special Publications.
- Guardiano, F.B., Srivastava, R.M., 1993. Multivariate geostatistics: beyond Bivariate moments. In: Soares, A. (Ed.), Geostatistics Tróia '92. Springer, Netherlands, pp. 133-144 number 5 in Quantitative Geology and Geostatistics.
- Howard, A.D., 1992. Modeling channel migration and floodplain sedimentation in meandering streams, Lowl, Floodplain Rivers Geomorphol, Perspect, 1-41,
- Howard, A.D., Hemberger, A.T., 1991. Multivariate characterization of meandering. Geomorphology 4, 161-186. http://dx.doi.org/10.1016/0169-555X(91)90002-R.
- Ikeda, S., Parker, G., Sawai, K., 1981. Bend theory of river meanders. Part 1. Linear development. J. Fluid Mech. 112, 363-377. http://dx.doi.org/10.1017/ S0022112081000451
- Imran, J., Islam, M.A., Kassem, A., 2008. The orientation of helical flow in curved

channels" by Corney et al. Sedimentology vol. 53, 249-257. http://dx.doi.org/ 10.1111/j.1365-3091.2007.00924.x discussion. Sedimentology 55, 235-239.

- Imran, J., Parker, G., Pirmez, C., 1999. A nonlinear model of flow in meandering submarine and subaerial channels. J. Fluid Mech. 400, 295-331. http:// dx.doi.org/10.1017/S0022112099006515.
- Janocko, M., Nemec, W., Henriksen, S., Warchoł, M., 2013. The diversity of deepwater sinuous channel belts and slope valley-fill complexes. Mar. Pet. Geol. 41, 7–34. http://dx.doi.org/10.1016/j.marpetgeo.2012.06.012.
- Labourdette, R., 2007. Integrated three-dimensional modeling approach of stacked turbidite channels. AAPG Bull. 91, 1603–1618. http://dx.doi.org/10.1306/ 06210706143
- Labourdette, R., 2008. 'LOSCS' Lateral Offset Stacked Channel Simulations: towards geometrical modelling of turbidite elementary channels. Basin Res. 20, 431–444. http://dx.doi.org/10.1111/j.1365-2117.2008.00361.x.
- Labourdette, R., Bez, M., 2010. Element migration in turbidite systems: random or systematic depositional processes? AAPG Bull. 94, 345-368. http://dx.doi.org/ 101306/09010909035
- Labourdette, R., Poncet, J., Seguin, J., Temple, F., Hegre, J., Irving, A., 2006. Threedimensional modelling of stacked turbidite channels in West Africa: impact on dynamic reservoir simulations. Pet. Geosci. 12, 335-345. http://dx.doi.org/ 10.1144/1354-079306-705
- Lindenmayer, A., 1968. Mathematical models for cellular interactions in development I. Filaments with one-sided inputs. J. Theor. Biol. 18, 280-299. http:// dx.doi.org/10.1016/0022-5193(68)90079-9.
- Lopez, S., 2003. Modélisation de réservoirs chenalisés méandriformes : une approche génétique et stochastique. Ph.D. thesis. Ecole Nationale Supérieure des Mines de Paris.
- Maier, K.L., Fildani, A., McHargue, T.R., Paull, C.K., Graham, S.A., Caress, D.W., 2012. Punctuated deep-water channel migration: high-resolution subsurface data from the Lucia Chica channel system, offshore California, U.S.A. J. Sediment. Res. 82, 1–8. http://dx.doi.org/10.2110/jsr.2012.10.
- Mariethoz, G., Comunian, A., Irarrazaval, I., Renard, P., 2014. Analog-based meandering channel simulation. Water Resour. Res. 50, 836-854. http:// dx.doi.org/10.1002/2013WR013730.
- Mariethoz, G., Renard, P., Straubhaar, J., 2010. The Direct Sampling method to perform multiple-point geostatistical simulations. Water Resour. Res. 46, W11536. http://dx.doi.org/10.1029/2008WR007621.
- Mayall, M., Jones, E., Casey, M., 2006. Turbidite channel reservoirs key elements in facies prediction and effective development. Mar. Pet. Geol. 23, 821-841. http:// dx.doi.org/10.1016/j.marpetgeo.2006.08.001.
- Mayall, M., O'Byrne, C., 2002. Reservoir Prediction and Development Challenges in Turbidite Slope Channels. Offshore Technology Conference, Houston, Texas U.S.A. http://dx.doi.org/10.4043/14029-MS.
- McHargue, T., Pyrcz, M., Sullivan, M., Clark, J., Fildani, A., Romans, B., Covault, J., Levy, M., Posamentier, H., Drinkwater, N., 2011. Architecture of turbidite channel systems on the continental slope: patterns and predictions. Mar. Pet. Geol. 28, 728-743. http://dx.doi.org/10.1016/j.marpetgeo.2010.07.008.
- Nakajima, T., Peakall, J., McCaffrey, W.D., Paton, D.A., Thompson, P.J.P., 2009. Outerbank bars: a new intra-channel architectural element within sinuous submarine slope channels. J. Sediment. Res. 79, 872-886. http://dx.doi.org/10.2110/ jsr.2009.094.
- O'Neill, M.P., Abrahams, A.D., 1986. Objective identification of meanders and bends. J. Hydrol. 83, 337-353. http://dx.doi.org/10.1016/0022-1694(86)90160-5
- Parker, G., Shimizu, Y., Wilkerson, G.V., Eke, E.C., Abad, J.D., Lauer, J.W., Paola, C., Dietrich, W.E., Voller, V.R., 2011. A new framework for modeling the migration of meandering rivers. Earth Surf. Process. Landforms 36, 70-86. http:// dx.doi.org/10.1002/esp.2113.
- Peakall, J., McCaffrey, B., Kneller, B., 2000. A process model for the evolution, morphology, and architecture of sinuous submarine channels. J. Sediment. Res. 434-448. http://dx.doi.org/10.1306/2DC4091C-0E47-11D7-8643000102C1865D
- Posamentier, H.W., 2003. Depositional elements associated with a basin floor channel-levee system: case study from the Gulf of Mexico. Mar. Pet. Geol. 20, 677-690. http://dx.doi.org/10.1016/j.marpetgeo.2003.01.002.
- Pyrcz, M., Boisvert, J., Deutsch, C., 2009. ALLUVSIM: a program for event-based stochastic modeling of fluvial depositional systems. Comput. Geosci. 35, 1671-1685. http://dx.doi.org/10.1016/j.cageo.2008.09.012.
- Pyrcz, M.J., Deutsch, C.V., 2005. Conditioning event-based fluvial models. In: Leuangthong, O., Deutsch, C.V. (Eds.), Geostatistics Banff 2004. Springer, Netherlands, pp. 135–144 number 14 in Quantitative Geology and Geostatistics.
- Pyrcz, M.J., Sech, R.P., Covault, J.A., Willis, B.J., Sylvester, Z., Sun, T., Garner, D., 2015. Stratigraphic rule-based reservoir modeling. Bull. Can. Pet. Geol. 63, 287-303. http://dx.doi.org/10.2113/gscpgbull.63.4.287
- Rongier, G., 2016. Connectivity of Channelized Sedimentary Bodies: Analysis and Simulation Strategies in Subsurface Modeling. Ph.D. thesis. Université de Lorraine/Université de Neuchâtel, Nancy, France/Neuchâtel, Switzerland. http://hal. archives-ouvertes.fr/tel-01371350.
- Rongier, G., Collon, P., Renard, P., Straubhaar, J., Sausse, J., 2016. Comparing connected structures in ensemble of random fields. Adv. Water Resour. 96, 145-169. http://dx.doi.org/10.1016/j.advwatres.2016.07.008.
- Rongier, G., Collon-Drouaillet, P., Filipponi, M., 2014. Simulation of 3D karst conduits with an object-distance based method integrating geological knowledge. Geomorphology 217, 152-164. http://dx.doi.org/10.1016/ j.geomorph.2014.04.024.
- Ruiu, J., Caumon, G., Viseur, S., 2015. Modeling channel forms and related

sedimentary objects using a boundary representation based on non-uniform rational B-Splines. Math. Geosci. 1–26. http://dx.doi.org/10.1007/s11004-015-9629-3.

- Schwenk, J., Lanzoni, S., Foufoula-Georgiou, E., 2015. The life of a meander bend: connecting shape and dynamics via analysis of a numerical model. J. Geophys.
- connecting shape and dynamics via analysis of a numerical model. J. Geophys.
  Res. Earth Surf. http://dx.doi.org/10.1002/2014JF003252, 2014JF003252.
  Sumner, E.J., Peakall, J., Dorrell, R.M., Parsons, D.R., Darby, S.E., Wynn, R.B., McPhail, S.D., Perrett, J., Webb, A., White, D., 2014. Driven around the bend: spatial evolution and controls on the orientation of helical bend flow in a natural submarine gravity current. J. Geophys. Res. Ocean. 119, 898–913. http://

dx.doi.org/10.1002/2013 C009008.

- Teles, V., de Marsily, G., Perrier, E., 1998. Sur une nouvelle approche de modélisation de la mise en place des sédiments dans une plaine alluviale pour en représenter l'hétérogénéité. Comptes Rendus de l'Académie des Sci. - Ser. IIA - Earth Planet. Sci. 327, 597–606. http://dx.doi.org/10.1016/S1251-8050(99)80113-X.
- Viseur, S., 2001. Simulation stochastique basée-objet de chenaux. Ph.D. thesis. Institut National Polytechnique de Lorraine, Vandoeuvre-lès-Nancy, France.
- Zhang, X., Pyrcz, M.J., Deutsch, C.V., 2009. Stochastic surface modeling of deepwater depositional systems for improved reservoir models. J. Pet. Sci. Eng. 68, 118–134. http://dx.doi.org/10.1016/j.petrol.2009.06.019.